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# Simulation of the interannual variations of aerosols in China: role of variations in meteorological parameters

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## Abstract

We used the nested grid version of the global three-dimensional Goddard Earth Observing System chemical transport model (GEOS-Chem) to examine the interannual variations (IAVs) of aerosols over heavily polluted regions in China for years 2004–2012. The role of variations in meteorological parameters was quantified by a simulation with fixed anthropogenic emissions at year 2006 levels and changes in meteorological parameters over 2004–2012. Simulated $PM_{2.5}$ (particles with a diameter of 2.5 µm or less) aerosol concentrations exhibited large IAVs in North China (NC, 32–42° N, 110–120° E), with regionally averaged absolute percent departure from the mean (APDM) values of 17, 14, 14, and 11 % in December-January-February (DJF), March-April-May (MAM), June-July-August (JJA), and September-October-November (SON), respectively. Over South China (SC, 22–32° N, 110–120° E), the IAVs in $PM_{2.5}$ were found to be the largest in JJA, with the regional mean APDM values of 14 % in JJA and of about 9 % in other seasons. Concentrations of $PM_{2.5}$ over the Sichuan Basin (SCB, 27–33° N, 102–110° E) were simulated to have the smallest IAVs among the polluted regions examined in this work, with the APDM values of 8–9 % in all seasons. All aerosol species (sulfate, nitrate, ammonium, black carbon, and organic carbon) were simulated to have the largest IAVs over NC in DJF, corresponding to the large variations in meteorological parameters over NC in this season. Process analyses were performed to identify the key meteorological parameters that determined the IAVs of different aerosol species in different regions. While the variations in temperature and specific humidity, which influenced the gas-phase formation of sulfate, jointly determined the IAVs of sulfate over NC in both DJF and JJA, wind (or convergence of wind) in DJF and precipitation in JJA were the dominant meteorological factors to influence IAVs of sulfate over SC and the SCB. The IAVs in temperature and specific humidity influenced gas-to-aerosol partitioning, which were the major factors that led to the IAVs of nitrate aerosol in China. The IAVs in wind and precipitation were found to drive the IAVs of organic carbon aerosol. We also compared the IAVs of aerosols simulated with





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variations in meteorological parameters alone with those simulated with variations in both meteorological parameters and anthropogenic emissions; the variations in meteorological fields were found to dominate the IAVs of aerosols in China.

## 1 Introduction

Aerosols are major air pollutants that have adverse effects on human health, reduce atmospheric visibility, and influence global climate change. With the rapid economic development in China over the past decades, concentrations of aerosols are now among the highest in the world (Fu et al., 2008; Cao et al., 2012; Sun et al., 2013) driven mainly by the increases in direct and precursor emissions (Streets et al., 2003). Aerosol concentrations in China have variations on different time scales (L. Zhang et al., 2010; Zhu et al., 2012); we aim to understand interannual variations (IAVs) of aerosols in this study.

The IAVs of aerosols were usually quantified in previous studies by statistical variables such as standard deviation (SD), relative standard deviation (RSD), mean absolute deviation (MAD), and absolute percent departure from the mean (APDM), which are defined as $\mathrm{SD} = \sqrt{\frac{1}{n}\sum_{i=1}^{n}\left(C_i - \frac{1}{n}\sum_{i=1}^{n}C_i\right)^2}$, $\mathrm{RSD} = 100\,\% \times \mathrm{SD} / \left(\frac{1}{n}\sum_{i=1}^{n}C_i\right)$, $\mathrm{MAD} = \frac{1}{n}\sum_{i=1}^{n}\left|C_i - \frac{1}{n}\sum_{i=1}^{n}C_i\right|$, and $\mathrm{APDM} = 100\,\% \times \mathrm{MAD} / \left(\frac{1}{n}\sum_{i=1}^{n}C_i\right)$, where $C_i$ is aerosol concentration of year $i$, and $n$ is the number of years examined. Therefore SD and MAD represent the absolute IAVs in aerosol concentration, and RSD and APDM represent the IAVs relative to the average concentration over the $n$ years.

Large IAVs of aerosols have been reported in previous studies for different aerosol species in different regions. Mahowald et al. (2003) showed that annual mean mineral dust concentrations measured at 10 sites in the United States over 1979–2000 had RSD values of 57–101 %. Habib et al. (2006) found by using the Total Ozone Mapping Spectrometer Absorbing Aerosol Index datasets that the absorbing aerosol column burdens averaged over April–May of 1981–1992 exhibited RSD of 16.0–30.3 % in different





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regions of India. Alston et al. (2012) showed by using ground-based measurements at 41 sites in the southeastern United States (29 sites of $PM_{2.5}$ measurements provided by Environmental Protection Agency and 12 sites by Georgia Dept. of Natural Resources) that monthly-mean $PM_{2.5}$ concentrations in that region had APDM values of 5–10 % over 2000–2009. They also showed by using three sets of aerosol optical depth (AOD) from Moderate Resolution Imaging Spectroradiometer (MODIS) Terra, Multi-angle Imaging SpectroRadiometer Terra, and MODIS Aqua that monthly-mean AOD in the southeastern United States had RSD values of 15–25 % over 2000–2009.

Observations and modeling studies have also shown that aerosols in China have large IAVs. Qu et al. (2010) reconstructed $PM_{10}$ aerosol concentrations at 86 Chinese cities using records of air pollution index from summer 2000 to winter 2006, and reported that seasonal-median $PM_{10}$ levels exhibited APDM values of 15–35 % in those cities. Yang et al. (2011) collected weekly samples of carbonaceous aerosols over 2005–2008 at two sites in Beijing and reported that year-to-year changes in emission and meteorology altered annual-average fine organic carbon (OC) concentrations at the rural site in Beijing by as much as 27.1 % over the observational time period.

The IAVs of aerosols are influenced by both emissions and meteorology. Meteorological parameters influence aerosol concentrations through altering emissions, chemical reactions, transport, and deposition. For example, increases in temperature enhance chemical production of sulfate in the atmosphere (Aw and Kleeman, 2003; Dawson et al., 2007; Kleeman, 2008) and decrease nitrate aerosol formation (Bellouin et al., 2011; Liao et al., 2006; Pye et al., 2009; Racherla and Adams, 2006). Aerosol concentrations decrease with increasing precipitation as wet deposition provides the main aerosol sink (Balkanski et al., 1993; Dawson et al., 2007), and changes in ventilation (wind speed, mixing depth) have large impacts on aerosols since aerosols are mainly influenced by local meteorological conditions. Tai et al. (2010) found that daily variation in meteorology as described by the multiple linear regression could explain up to 50 % of $PM_{2.5}$ variability based on an 11 year (1998–2008) observational record over





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the contiguous United States, with temperature, relative humidity, precipitation, and circulation all being important predictors.

Since concentrations of chemical species are the net results from comprehensive physical and chemical processes, the Integrated Process Rates (IPR) (Im et al., 2011) have been used to identify the dominant processes (such as horizontal and vertical transport, emissions of primary species, gas-phase chemistry, dry deposition, cloud processes, and aerosol processes) that influence the concentrations of chemical species in the Community Multi-scale Air Quality model (CMAQ) for episodic events (Jose et al., 2002; Goncalves et al., 2009) as well as for yearly (Zhang et al., 2009) to decadal simulations (Civerolo et al., 2010). The IPR analyses in these studies ranked the roles of different processes in the formation and fate of a chemical species. For example, the horizontal flows and gas-phase chemical reactions in the morning and the vertical flows in the afternoon were found to be the main factors in the formation of surface $O_3$ during a photochemical pollution episode in the coastal area of South-Western Europe in summer (Goncalves et al., 2009). Recently, a similar process analysis scheme was implemented within the Weather Research and Forecasting model coupled with Chemistry (WRF Chem) to understand the key photochemical and physical processes for the formation of $O_3$ (Jiang et al., 2012) and $PM_{10}$ (Jiang et al., 2013).

The scientific goals of this work are: (1) to quantify the IAVs in surface-layer aerosol concentrations in China resulted from the variations in meteorological conditions during 2004–2012, using the global three-dimensional chemical transport model GEOS-Chem, and (2) to identify the key meteorological parameters that influenced the IAVs of aerosols in different polluted regions of China by the IPR analyses. Section 2 describes the model, emissions, and numerical experiments. Section 3 presents simulated distributions of aerosols and IAVs in concentrations of different aerosol species in China averaged over 2004–2012. The key meteorological parameters that influenced IAVs of aerosols are examined by IPR in Sect. 4. Section 5 discusses the impacts of anthropogenic and natural emissions on IAVs of aerosols in China.





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## 2 Model description and numerical experiments

### 2.1 GEOS-Chem Model

We simulated aerosols using the global chemical transport model GEOS-Chem (version 9-01-02) driven by the GEOS-5 assimilated meteorological fields from the Goddard Earth Observing System of the NASA Global Modeling and Assimilation Office. We used the nested-grid capability of the GEOS-Chem model over East Asia (11° S–55° N, 70–150° E) with a horizontal resolution of 0.5° latitude by 0.667° longitude and 47 vertical layers up to 0.01 hPa (Chen et al., 2009). Chemical boundary conditions were from the global simulations performed at 4° × 5° horizontal resolution.

The GEOS-Chem model has fully coupled $O_3$–$NO_x$–hydrocarbon chemistry and aerosols including sulfate ($SO_4^{2-}$), nitrate ($NO_3^-$), ammonium ($NH_4^+$) (Park et al., 2004; Pye et al., 2009), OC and BC (Park et al., 2003), sea salt (Alexander et al., 2005), and mineral dust (Fairlie et al., 2007). The gas-aerosol partitioning of nitric acid and ammonia is calculated using the ISORROPIA II thermodynamic equilibrium model (Fountoukis and Nenes, 2007). Wet deposition of soluble aerosols and gases follows the scheme of Liu et al. (2001) and dry deposition follows a standard resistance-in-series model of Wesely (1989). We do not examine IAVs of mineral dust and sea salt aerosols in this study, because sea salt aerosol is not a major aerosol component in China based on measurements (Ye et al., 2003; Duan et al., 2006) and mineral dust aerosol simulation has very large uncertainties (Fairlie et al., 2007, 2010).

Considering the large uncertainties in chemistry schemes of secondary organic aerosol (SOA), SOA in our simulations was assumed to be the 10 % carbon yield of OC from biogenic terpenes (Park et al., 2003) and 2 % carbon yield of OC from biogenic isoprene (van Donkelaar et al., 2007).





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## 2.2 Emissions

Global emissions of aerosol precursors and aerosols in the GEOS-Chem model generally follow Park et al. (2003) and Park et al. (2004), in which anthropogenic emissions of $NO_x$, $SO_2$, BC and OC (including emissions from power, industry, residential, and transportation) in the Asian domain have been overwritten by David Streets' 2006 emission inventory (http://mic.greenresource.cn/intex-b2006). Estimates of $NH_3$ emissions in China showed large uncertainties in previous studies (Streets et al., 2003; Kim et al., 2006; Y. Zhang et al., 2010; Huang et al., 2011, 2012). We used in our simulations the most recent estimate of $NH_3$ emissions in China by Huang et al. (2012), which was 9.8 Tg $yr^{-1}$, instead of 13.5 Tg $yr^{-1}$ from Streets et al. (2003). Table 1 summarizes year 2006 annual emissions of $NO_x$, $SO_2$, $NH_3$, OC, and BC in Asia (70–150° E, 11° S–55° N) and eastern China (98–130° E, 20–55° N).

Natural $NO_x$ emissions from lightning and soil were described by Sauvage et al. (2007) and Yienger and Levy (1995). Natural $NH_3$ emissions from soil, vegetation, and the oceans were from the Global Emissions Inventory Activity inventory (Bouwman et al., 1997). Biomass burning emissions were from the monthly Global Fire Emissions Database-v2 inventory (van der Werf et al., 2006). Biogenic VOC emissions were calculated from Model of Emissions of Gases and Aerosols from Nature (Guenther et al., 2006).

## 2.3 Numerical experiments

To quantify IAVs of aerosols over 2004–2012, we performed the following simulations of aerosols in China using the GEOS-Chem model driven by the GEOS-5 meteorological fields:

1. ANNmet: the simulation to examine how the IAVs of aerosols were influenced by variations in meteorological parameters. Meteorological fields and biomass





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burning emissions were allowed to vary from 2004 to 2012, while anthropogenic emissions were kept at the year 2006 values.

2. ANNmet_ATM: sensitivity simulation for 2004–2012 to examine the sensitivity of IAVs of aerosols to variations in atmospheric conditions alone. All natural emissions (such as soil $NO_x$, lightning $NO_x$ as well as biogenic sources) that were sensitive to meteorological parameters were turned off. Anthropogenic emissions were kept at the year 2006 values. Meteorological fields and biomass burning emissions were allowed to vary from 2004 to 2012.

3. ANNall: simulation of aerosols for years 2004–2012 with yearly varying meteorological parameters, biomass, natural and anthropogenic emissions. The IAVs in anthropogenic emissions over 2004–2012 were obtained by using scaling factors; the annual scaling factors for $NO_x$ were taken from Zhang et al. (2012) and those for $SO_2$, OC, and BC were taken from Lu et al. (2011).

In ANNmet, the IAVs in meteorological fields influenced aerosol concentrations in two ways. First, changes in meteorological parameters influenced chemical reactions, transport, and deposition of aerosols. Second, precursor emissions from natural sources varied with meteorological fields. We further separated these effects by performing one sensitivity simulation ANNmet_ ATM with natural emissions turned off. Aerosol concentrations from ANNall will be used for understanding the role of anthropogenic emissions in IAVs of aerosols, as aerosol concentrations from ANNall are compared with those from ANNmet.

The presentations of the IAVs of aerosols will be focused on three polluted regions in China, North China (NC, 32–42° N, 110–120° E), South China (SC, 22–32° N, 110–120° E), and the Sichuan Basin (SCB, 27–33° N, 102–110° E), as defined in Fig. 1.





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## 3 Simulated IAVs of aerosols resulted from IAVs of meteorological parameters alone

### 3.1 Simulated distributions of aerosol concentrations

Figure 2 shows simulated seasonal-mean surface-layer concentrations of sulfate, nitrate, ammonium, OC, BC, and $PM_{2.5}$ (sum of sulfate, nitrate, ammonium, OC, and BC) averaged over 2004–2012 of simulation ANNmet. Simulated aerosol concentrations were high over polluted eastern China throughout the year. Sulfate concentrations in NC were 15–20 $\mu g\,m^{-3}$ in June-July-August (JJA) with the strong photochemistry in that season, and the concentrations over SC showed small values of 3–10 $\mu g\,m^{-3}$ in JJA as a result of the large precipitation associated with the summer monsoon. The maximum sulfate concentrations of 25–30 $\mu g\,m^{-3}$ were simulated over SCB in December-January-February (DJF), as a result of the large $SO_2$ emissions from winter heating.

Simulated nitrate concentrations over NC were in the range of 15–40 $\mu g\,m^{-3}$ throughout the year, with maximum concentrations of about 40 $\mu g\,m^{-3}$ in DJF. High $NO_x$ emissions and low temperatures favored nitrate formation in DJF. Nitrate concentrations showed values of 5–15 µg over SC and SCB $m^{-3}$ in JJA when temperatures were the highest. Simulated ammonium concentrations were in the range of 5–20 $\mu g\,m^{-3}$ over NC, SC, and SCB, with seasonal variations in these regions following those of nitrate.

The simulated distributions of OC and BC were similar to those of their emissions. Simulated OC and BC concentrations were high in DJF and September-October-November (SON) and low in March-April-May (MAM) and JJA, owing to the seasonal variations in precipitation.

Simulated $PM_{2.5}$ concentrations were in the range of 70–90 $\mu g\,m^{-3}$ in NC throughout the year, which were generally higher than the concentrations in SC and SCB. In the surface layer, nitrate was predicted to be the most abundant aerosol species over eastern China, followed by sulfate, ammonium, OC, and BC. Wang et al. (2013) reported that high nitrate in the GEOS-Chem model is likely caused by the overestimate of $NH_3$ emissions.





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### 3.2 Simulated IAVs of aerosols

Figures 3 and 4 show, respectively, the MAD and APDM values of seasonal mean surface-layer concentrations of sulfate, nitrate, ammonium, OC, BC, and $PM_{2.5}$ simulated in ANNmet. We also present the domain-averaged values of MAD and APDM for NC, SC, and SCB in Tables 2 and 3.

The MAD and APDM values of sulfate, nitrate, and ammonium from ANNmet indicated that concentrations of these species in NC had larger IAVs than those in SC. Over NC, SC, and SCB, sulfate aerosol showed regional mean MAD values of 0.9–2.0 $\mu g\,m^{-3}$ and APDM values of 10–24 % (Tables 2 and 3). The largest IAVs of sulfate were found over NC in DJF, with MAD values of exceeding 2.5 $\mu g\,m^{-3}$ and APDM values of exceeding 24 %. These IAVs were significant, considering that these were averages over 2004–2012; year by year variations were larger than the averages reported here. Our simulated IAV of sulfate was close to the IAV of 14–20 % reported by Gong et al. (2010) for sulfate at the Canadian High Arctic.

The MAD and APDM values obtained in simulation ANNmet showed that nitrate concentrations also had large IAVs. Over NC where nitrate concentrations were the highest, MAD and APDM values were, 1.8–2.4 $\mu g\,m^{-3}$ and 13–18 %, respectively. The APDM values of nitrate in NC did not show large variations with season, which were larger than the APDM values of sulfate in JJA but smaller than those of sulfate in DJF. The distribution and magnitude of IAVs of ammonium generally followed those of nitrate over polluted eastern China.

The spatial pattern of either MAD or APDM of OC was similar to that of BC in all seasons. OC exhibited seasonal mean MAD values of 0.2–0.5 $\mu g\,m^{-3}$ and APDM values of 6–12 % in polluted NC, SC, and SCB throughout the year. BC is a chemically inert tracer that responds to changes in meteorological parameters. The APDM values of BC were about the same as those of OC, except that the APDM values of BC were smaller in NC in MAM and in SC in DJF, MAM, and JJA.





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As a result of the high ratios of nitrate to $PM_{2.5}$ concentrations, the patterns of the IAVs of $PM_{2.5}$ were similar to those of nitrate. The IAVs of $PM_{2.5}$ concentrations were the largest in NC; the regional mean MAD values were in the range of 4.4–6.1 $\mu g\,m^{-3}$ and the APDM values were 17, 14, 14, and 11 % in DJF, MAM, JJA, and SON (Tables 2 and 3), respectively. Over SC, MAD and APDM values were 2.8–3.7 $\mu g\,m^{-3}$ and 9–14 %, respectively; the maximum APDM value of 14 % was found in JJA and the highest APDM values were about 9 % in other seasons (Tables 2 and 3). $PM_{2.5}$ showed the smallest IAVs over SCB, with MAD and APDM values of, respectively, 2.0–3.1 $\mu g\,m^{-3}$ and 8–9 % in all seasons.

### 3.3 Comparisons of simulated IAVs of aerosols with measurements

Simulated concentrations of sulfate, nitrate, and ammonium aerosols in China have been evaluated in the study of Wang et al. (2013) and those of carbonaceous aerosols in China have been evaluated by Fu et al. (2012), both of which used the same one-way nested-grid capability of the GEOS-Chem. Wang et al. (2013) found that simulated concentrations of sulfate, nitrate and ammonium at 22 sites in East Asia exhibited annual mean biases of −10, +31, and +35 %, respectively, and Fu et al. (2012) showed that the simulated annual mean concentrations of BC and OC averaged over rural and background sites were underestimated by 56 and 76 %, respectively.

For the purpose of this study, we evaluated the model's performance in simulating the IAVs of aerosols by comparing simulated aerosol optical depths (AODs) with satellite measurements, because of the lack of long-term ground-based measurements in China (Chan and Yao, 2008). The Level 3 MODIS/Terra monthly products (MOD08_M3, http://ladsweb.nascom.nasa.gov/) with $1^\circ \times 1^\circ$ equal-angle global grid were obtained from NASA LAADS (Level 1 and Atmosphere Archive and Distribution System). Collections 5 and 5.1 contain the time series of AODs from March 2000 to the present. We used AODs at the 550 nm wavelength, which incorporated only the highest quality retrievals. Over East Asia, the MODIS measurements have been well validated through





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many studies (Chin et al., 2004; Park et al., 2011). In the GEOS-Chem model, the AODs of sulfate, nitrate, ammonium, OC, BC, sea salt, and dust aerosols were calculated based on aerosol mass concentration, extinction efficiency, effective radius, particle mass density, and the assumed aerosol size distribution (Drury et al., 2010). The hygroscopic growth of each aerosol species with relative humidity was accounted for, using the hygroscopic growth factors listed in Martin et al. (2003).

We compared simulated and observed IAVs in AODs for the cities of Beijing (39.5° N, 116.2° E), Changsha (28.1° N, 112.6° E) and Chengdu (30.7° N, 104° E) (Fig. 5), which were chosen to represent the model performance in NC, SC and SCB regions, respectively. The correlation coefficients between observed and modeled monthly mean AODs were 0.83, 0.34, 0.13 for Beijing, Changsha, and Chengdu, respectively. The large correlation coefficient for Beijing indicates that the model was able to capture to some extent the observed IAVs in NC. The small correlation coefficients in Changsha and Chengdu can be explained in part by the complex topography and satellite limitations like cloud contamination (Xia et al., 2004). Note that in simulation ANNmet, simulated AODs correlated well with the simulated column burdens and surface-layer concentrations of $PM_{2.5}$; the correlation coefficients between simulated monthly mean AODs and column burdens (surface-layer concentrations) of $PM_{2.5}$ were 0.95 (0.81), 0.94 (0.81), and 0.91 (0.77) over Beijing, Changsha, and Chengdu, respectively. The IAVs of observed AODs agreed fairly well with the IAVs of surface-layer aerosol concentrations. For example, the seasonal-mean APDM values of observed AODs were 18.4 % (DJF), 14.9 % (MAM), 24.2 % (JJA), and 16.3 % (SON) for Beijing, close to the seasonal-mean APDM values of surface-layer $PM_{2.5}$ shown in Fig. 4.





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## 4 Understanding the IAVs of aerosols by process analyses

### 4.1 IAVs of meteorological parameters

Figure 6 shows seasonal mean temperature, specific humidity, precipitation, 850 hPa zonal and meridional winds for DJF and JJA. All these meteorological fields were averaged over years 2004–2012. Over central and eastern China, temperature, specific humidity and precipitation generally exhibited much larger values in JJA than in DJF. At 850 hPa, strong westerlies were found in NC in DJF, and prevailing southerlies occurred in JJA, reflecting the typical features of winds in China.

Figures 7 and 8 show, respectively, the MAD and APDM values of surface-layer temperature, specific humidity, and precipitation for DJF and JJA. Over NC, SC, and SCB, temperature exhibited MAD values of 1.00–2.00 K in DJF and of 0–1.25 K in JJA. The APDM values of temperature in DJF were generally larger than those in JJA. Piao et al. (2003) also showed that the largest IAV of temperature in eastern China was in winter (December and February) based on the reanalyzed temperatures over 1982–1999. Specific humidity showed APDM values of 6–20 % in DJF and of 2–8 % in JJA over central and eastern China, although the MAD values of specific humidity over NC and SCB were larger in JJA than in DJF. For precipitation, APDM values of 20–80 % and 10–30 % were calculated in DJF and JJA, respectively, and the APDM values of precipitation were larger in NC than in SC, which agreed with those reported in Qian and Lin (2005).

### 4.2 Process analyses

The concentrations of aerosols are determined by emissions, chemical reactions, transport, and deposition. Therefore, the IAVs of aerosols are influenced by IAV of each of these processes. The weighted contribution of each process to IAV of an aerosol is estimated here using $\%PC_i = MAD_i / \sum_i^n MAD_i$, where $n$ is the number of processes considered, $MAD_i$ is the MAD value of process $i$ and $\%PC_i$ is the relative contribution





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of process *i* to the sum of the contributions from all processes (Im et al., 2011). Once the most important processes are selected from this approach, meteorological variables to which the processes are sensitive to are classified as the key meteorological parameters that lead to IAV of the aerosol. Since we aimed to examine the IAVs in surface-layer aerosol concentrations, our process analyses for an aerosol species were performed for each region (NC, SC, or SCB) from the surface to 1 km altitude.

#### 4.2.1 Sulfate

Processes that influence the IAVs of sulfate concentrations include anthropogenic emissions, formation pathways (gas-phase oxidation of $SO_2$ by OH, and in-cloud oxidation of $SO_2$ by ozone and hydrogen peroxide), mass fluxes through the four lateral boundaries and the upper boundary at 1 km, wet and dry deposition. Figure 9 shows the mass fluxes and APDM values and Fig. 10 shows the MAD values and relative contributions of individual atmospheric processes to sulfate concentrations in different regions for DJF and JJA. Transport fluxes were calculated at the boundaries while other sources and sinks were the sums over the grid cells of the region under 1 km.

With respect to sulfate budget over NC in DJF, vertical flux through the top side and gas-phase sulfate formation by reaction of $SO_2$ with OH had the largest values of 0.041 Tg S season$^{-1}$ and 0.019 Tg S season$^{-1}$, with APDM values of 17.2 and 65.0 %, respectively (Fig. 9a). Among the four horizontal fluxes through the lateral boundaries of NC, the flux through the south boundary had the largest value of 0.017 Tg S season$^{-1}$ with an APDM value of 29.1 %, while that through the west boundary had the smallest value of 0.003 Tg S season$^{-1}$ with a large APDM value of 62.1 %. The in-cloud reactions of $SO_2$ with $O_3$ and $H_2O_2$ contributed relatively small to sulfate formation, by 0.005 Tg S season$^{-1}$ and 0.004 Tg S season$^{-1}$, respectively, but both of which had large APDM values of about 53 %. Wet deposition was 0.006 Tg S season$^{-1}$ with the highest APDM value of 62.1 %. Dry deposition was the smallest in terms of both flux and APDM value. Figure 10 shows that gas-phase oxidation of $SO_2$ by OH was the most important





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process that contributed to the IAV of sulfate, with MAD value of 0.012 Tg S season$^{-1}$ and relative contribution of 30.1 %, followed by the vertical flux through the top side (0.007 Tg S season$^{-1}$, 17.8 %). The relative contributions of horizontal fluxes through the lateral boundaries were in the range of 3.4–11.3 %. The aqueous-phase oxidations of $SO_2$ by $O_3$ and $H_2O_2$ accounted for 5.7 and 6.2 % of the IAV of sulfate, respectively. The relative contribution by wet deposition was 8.4 % and that by dry deposition was 0.4 % (Fig. 10a). In JJA, gas-phase formation of sulfate increased to 0.039 Tg S season$^{-1}$ with relative contribution of 56.9 %, making it the only process that determined the IAV of sulfate concentration over NC in this season. Other processes were either small in flux or small in APDM value, leading to small contributions to IAV of sulfate in NC in JJA. Since gas-phase oxidation of $SO_2$ is sensitive to temperature and humidity (Yao et al., 2002; Zhang et al., 2012), we conclude that the variations in temperature and specific humidity were the key factors that drove the IAV of sulfate over NC. The relatively high MAD values of temperature in NC in both DJF and JJA (Fig. 7) as well as the large MAD values of specific humidity over NC in JJA (Fig. 7) support the above conclusion.

Similar analyses were performed for sulfate in SC. In DJF, the gas-phase formation of sulfate and wet deposition were the dominant source and sink of sulfate in this region (Fig. 9b). The vertical flux, flux through the south boundary, and wet deposition had the largest contributions to IAV of sulfate with relative contributions of 23.8, 17.6, and 15.3 %, respectively, indicating that wind and precipitation were the main meteorological factors that determined the IAV of sulfate concentration over SC in DJF. In JJA, sulfate formation from the reaction of $SO_2$ with OH had a large value of 0.030 Tg S season$^{-1}$, but this process had a very low APDM value of 6.4 % (Fig. 9b). Figure 10b showed that wet deposition and in-cloud oxidation of $SO_2$ by $H_2O_2$ were the prevailing processes that contributed, respectively, 19.7 and 17.9 % to the IAV of sulfate concentration in SC. These two processes corresponded well with the high MAD values of precipitation in JJA as shown in Fig. 7.





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With respect to sulfate budget over the SCB in DJF, vertical flux through the top side, wet deposition, and gas-phase formation of sulfate had the largest values of 0.050, 0.023, and 0.019 Tg S season$^{-1}$, respectively (Fig. 9c). The vertical flux through the top side was the most important process that contributed to the IAV of sulfate, with MAD value of 0.008 Tg S season$^{-1}$ and relative contribution of 32.9 %, followed by wet deposition (0.005 Tg S season$^{-1}$, 23.4 %) and in-cloud oxidation of $SO_2$ by $O_3$ (0.003 Tg S season$^{-1}$, 12.0 %) (Fig. 10c). We can infer from these analyses that wind (or convergence of winds) was the dominant meteorological factor to influence IAV of sulfate over SCB in DJF. In JJA, wet deposition was the largest contributor to the IAV of sulfate concentration with relative contribution of 35.0 %, which corresponded well with the high MAD values of precipitation in JJA (Fig. 7).

#### 4.2.2 Nitrate

The processes that influence nitrate concentration include gas-to-aerosol conversion of $HNO_3$ to form nitrate, mass fluxes through the four lateral boundaries and the upper boundary at 1 km, wet and dry deposition. Among all the processes, gas-to-aerosol conversion was found to be the key process that drove the IAV of nitrate over the three studied regions. In DJF (JJA), gas-to-aerosol conversion was calculated to have relative contributions of 56.3 % (70.6 %), 45.0 % (72.6 %), and 62.3 % (72.5 %) in NC, SC, and SCB, respectively (Fig. 12), indicating that the IAVs in temperature and specific humidity drove the IAVs of nitrate in these polluted regions. As reported by Dawson et al. (2007), temperature and humidity have the largest influences on gas-to-aerosol partitioning of $HNO_3$.

#### 4.2.3 Organic carbon

Processes that influence IAV of OC include mass fluxes through the four lateral boundaries and the upper boundary, emissions from anthropogenic, biomass, biofuel and biogenic sources, as well as wet and dry deposition. With respect to OC budget over





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NC, emissions had the largest mass fluxes in OC budget in both DJF and JJA (Fig. 13). However, Fig. 14 shows that transport fluxes were the most important processes to drive the IAV of OC in both seasons. The relative contributions of transport fluxes were in the range of 5.2–22.6 % in DJF and 7.9–25.0 % in JJA. Wet deposition made contributions of 8.5 % in DJF and 11.2 % in JJA. Note that the relative contributions of biogenic and biomass emissions were very small.

Similar analysis for SC showed that biofuel emission, fluxes through the north and south boundaries had the largest values in OC budget in DJF (Fig. 13). Vertical transport, transport through the south boundary, and flux through the north boundary had the largest contributions to the IAV of OC in DJF, with relative contributions of 23.9 %, 20.2 %, and 15.4 %, respectively (Fig. 14). In JJA, transport through the north boundary had the largest relative contribution of 23.0 %, followed by wet deposition of 16.3 %.

Over SCB in DJF, vertical transport had the largest relative contribution of 31.1 %, followed by mass flux through the south boundary of 24.1 % (Fig. 14). In JJA, wet deposition, vertical transport, and transport at the south boundary had the largest contributions to the IAV of OC, with relative contributions of 29.9 %, 15.3 %, and 11.4 %, respectively.

To conclude, wind was the most important meteorological parameter that drove the IAV of OC in the three regions. Precipitation also played a crucial role in JJA over SC and SCB.

## 5 Impacts of anthropogenic emissions and meteorology-sensitive natural emissions on IAVs of aerosols

Tables 2 and 3 show, respectively, the simulated MAD and APDM values of $SO_4^{2-}$, $NO_3^-$, $NH_4^+$, OC, BC, and $PM_{2.5}$ in ANNmet, ANNmet_ATM and ANNall experiments. As described in Sect. 2, the differences in APDM values between ANNmet and ANNall represented the IAVs of aerosols caused by variations in anthropogenic emissions, which were compared with the APDM values obtained in ANNmet to see the relative





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importance of anthropogenic emissions and meteorological parameters in the IAVs of aerosols. For sulfate (OC), the maximum difference in APDM value between ANNall and ANNmet was 4.2 % for sulfate in JJA in SCB (1.6 % for OC aerosol in MAM in SCB), which was relatively small as compared with the APDM value of 10.8 % (6.6 %) in ANNmet in that season and area. Nitrate, ammonium, and $PM_{2.5}$ exhibited larger differences in APDM values between ANNall and ANNmet, especially over SC and SCB in MAM, JJA, and SON. These results can be explained by that the IAVs of anthropogenic emissions of $NO_x$ and BC were larger than those of $SO_2$ and OC during 2004–2010 (Lu et al., 2011; Zhang et al., 2012). Note that for all aerosol species listed in Table 3, the differences in APDM values between ANNall and ANNmet were smaller than the APDM values in ANNmet, indicating that the variations in meteorological parameters played more important roles than variations in anthropogenic emissions in driving the IAVs of aerosols in China.

The differences in MAD and APDM values between ANNmet_ATM and ANNmet represented the IAVs of aerosols caused by variations in meteorology-sensitive natural emissions. The roles of natural emissions were generally small, except that the differences in APDM between ANNmet_ATM and ANNmet were large in DJF and MAM for OC over SC (Table 3), which were caused by the high biogenic emissions in the region (Fu and Liao, 2012).

## 6 Conclusions

We used the nested grid version of the GEOS-Chem model to estimate the role of meteorology in the IAVs of aerosols over China for years 2004–2012. We performed simulations ANNall (combined effects of variations in meteorology and anthropogenic emissions), ANNmet (effects of variations in meteorology alone), and ANNmet_ATM (same as ANNmet except that meteorology-sensitive natural emissions were turned off) to identify the key parameters that influence the IAVs of aerosols.





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We defined two parameters, mean absolute deviation (MAD) and absolute percent departure from the mean (APDM), to quantify the IAVs in concentrations of aerosols over 2004–2012. Results from simulation ANNmet showed that, driven by changes in meteorological parameters alone, the regional mean APDM values of sulfate, nitrate, and OC aerosols were in the ranges of 10.1–23.6 %, 9.3–18.2 %, and 5.6–12.1 %, respectively, over the studied regions (NC, SC, and SCB) throughout the year. As a result of the IAVs of individual aerosol species, simulated $PM_{2.5}$ aerosol concentrations exhibited large IAVs in NC, with regionally averaged APDM values of 17, 14, 14, and 11 % in DJF, MAM, JJA, and SON, respectively. Over SC, the IAVs in $PM_{2.5}$ were found to be the largest in JJA; the regional mean APDM value was 14 % in JJA and about 9 % in other seasons. Concentrations of $PM_{2.5}$ over SCB were simulated to have the smallest IAVs among the polluted regions examined in this work, with the APDM values of 8–9 % in all seasons. All aerosol species (sulfate, nitrate, ammonium, black carbon, and organic carbon) were simulated to have the largest IAVs over NC in DJF, corresponding to the large variations in meteorological parameters over NC in DJF.

We applied process analyses for sulfate, nitrate and OC to identify key meteorological parameters that led to IAVs of these aerosols over 2004–2012. For sulfate in NC, gas-phase formation of sulfate was found to be the key process that drove the IAV of sulfate in both DJF and JJA, with relative contribution of 30.1 % in DJF and of 56.9 % in JJA, inferring that the variations in temperature and specific humidity jointly determined the IAV of sulfate in NC. Over SC and SCB, the most important process that dominated IAV of sulfate in DJF was found to be the vertical flux through the top side with relative contribution of 23.8 % in SC and of 32.9 % in SCB, and the key process in JJA was found to be wet deposition with relative contributions of 19.7 and 35.0 % in SC and SCB, respectively. For nitrate, gas-to-aerosol conversion was found to be the key process that dominated the IAVs of nitrate over the three regions, with very high relative contributions of 45.0–62.3 % in DJF and 70.6–72.6 % in JJA, indicating that temperature and specific humidity were the major factors that drove the IAV of nitrate in China. For OC, transport was the most important process that influenced the IAV of





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OC throughout the year, and precipitation also played a crucial role in JJA over SC and SCB, associated with the East Asian summer monsoon precipitation.

We also examined the relative importance of anthropogenic emissions and meteorological parameters in the IAVs of aerosols. For all aerosol species (sulfate, nitrate, ammonium, BC, and OC), the differences in APDM values between ANNall and ANNmet were smaller than the APDM values in ANNmet, indicating that the variations in meteorological parameters played more important roles than variations in anthropogenic emissions in driving the IAVs of aerosols in China. The roles of natural emissions were generally small, except that the differences in APDM between ANNmet_ATM and ANNmet were large in DJF and MAM for OC over SC, which were caused by the high biogenic emissions in the region.

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**Table 1.** Summary of annual emissions of aerosols and aerosol precursors in Asia (70–150° E, 11° S–55° N) and eastern China (98–130° E, 20–55° N).

| Species | Asia | Eastern China |
|---|---|---|
| $NO_x$ ($Tg N yr^{-1}$) | | |
| Aircraft | 0.08 | 0.02 |
| Anthropogenic | 10.6 | 6.38 |
| Biomass burning | 1.09 | 0.14 |
| Biofuel | 0.02 | < 0.01 |
| Fertilizer | 0.35 | 0.12 |
| Lightning | 1.16 | 0.28 |
| Soil | 0.88 | 0.27 |
| Total | 14.18 | 7.21 |
| $SO_2$ ($Tg S yr^{-1}$) | | |
| Aircraft | 0.01 | < 0.01 |
| Anthropogenic | 23.85 | 15.98 |
| Biomass burning | 0.36 | 0.04 |
| Biofuel | < 0.01 | < 0.01 |
| Volcanoes | 3.99 | 0.07 |
| Ship | 0.48 | 0.05 |
| Total | 28.69 | 16.14 |
| $NH_3$ ($Tg N yr^{-1}$) | | |
| Anthropogenic | 14.18 | 7.26 |
| Natural | 2.33 | 0.49 |
| Biomass burning | 0.82 | 0.11 |
| Biofuel | 0.8 | 0.3 |
| Total | 18.13 | 8.16 |
| OC ($Tg C yr^{-1}$) | | |
| Anthropogenic | 1.54 | 1.03 |
| Biomass burning | 4.62 | 0.67 |
| Biofuel | 3.47 | 1.44 |
| Biogenic | 2.5 | 0.55 |
| Total | 12.13 | 3.69 |
| BC ($Tg C yr^{-1}$) | | |
| Anthropogenic | 1.52 | 0.96 |
| Biomass burning | 0.55 | 0.05 |
| Biofuel | 0.92 | 0.39 |
| Total | 2.99 | 1.4 |





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**Table 2.** Simulated MAD values ($\mu g\,m^{-3}$) of aerosols in ANNmet, ANNmet_ATM, and ANNall. For each season, the numbers in the brackets of column ANNmet_ATM (ANNall) were calculated as the differences in MAD values between ANNmet_ATM (ANNall) and ANNmet (ANNmet_ATM (ANNall) – ANNmet).

| Species | DJF | | | MAM | | | JJA | | | SON | | |
|---|---|---|---|---|---|---|---|---|---|---|---|---|
| | ANNmet | ANNmet_ATM | ANNall | ANNmet | ANNmet_ATM | ANNall | ANNmet | ANNmet_ATM | ANNall | ANNmet | ANNmet_ATM | ANNall |
| | | | | | | NC | | | | | | |
| $SO_4^{2-}$ | 2.0 | 2.1 (+0.1) | 2.0 (0) | 1.3 | 1.3 (0) | 1.3 (0) | 1.4 | 1.4 (0) | 1.5 (+0.1) | 1.5 | 1.5 (0) | 1.6 (+0.1) |
| $NO_3^-$ | 2.3 | 2.3 (0) | 2.6 (+0.3) | 1.8 | 1.8 (0) | 2.5 (+0.7) | 2.4 | 2.3 (−0.1) | 3.0 (+0.6) | 2.1 | 2.0 (−0.1) | 2.3 (+0.2) |
| $NH_4^+$ | 1.2 | 1.2 (0) | 1.2 (0) | 1.0 | 0.9 (−0.1) | 1.1 (+0.1) | 1.1 | 1.1 (0) | 1.3 (+0.2) | 1.0 | 1.0 (0) | 1.0 (0) |
| OC | 0.5 | 0.5 (0) | 0.6 (+0.1) | 0.3 | 0.3 (0) | 0.3 (0) | 0.3 | 0.3 (0) | 0.3 (0) | 0.3 | 0.3 (0) | 0.3 (0) |
| BC | 0.3 | 0.3 (0) | 0.3 (0) | 0.1 | 0.1 (0) | 0.2 (+0.1) | 0.1 | 0.1 (0) | 0.2 (+0.1) | 0.2 | 0.2 (0) | 0.2 (0) |
| $PM_{2.5}$ | 6.1 | 6.2 (+0.1) | 6.4 (+0.3) | 4.4 | 4.2 (−0.2) | 5.1 (+0.7) | 5.0 | 4.9 (−0.1) | 5.7 (+0.7) | 4.9 | 4.8 (−0.1) | 4.8 (−0.1) |
| | | | | | | SC | | | | | | |
| $SO_4^{2-}$ | 1.7 | 1.8 (+0.1) | 1.7 (0) | 1.2 | 1.2 (0) | 1.2 (0) | 0.9 | 0.9 (0) | 0.9 (0) | 1.1 | 1.1 (0) | 1.2 (+0.1) |
| $NO_3^-$ | 1.4 | 1.4 (0) | 1.7 (+0.3) | 1.2 | 1.2 (0) | 2.0 (+0.8) | 1.2 | 1.2 (0) | 1.8 (+0.6) | 1.3 | 1.3 (0) | 1.9 (+0.6) |
| $NH_4^+$ | 0.8 | 0.8 (0) | 0.8 (0) | 0.7 | 0.7 (0) | 0.9 (+0.2) | 0.6 | 0.6 (0) | 0.8 (+0.2) | 0.7 | 0.7 (0) | 0.8 (+0.1) |
| OC | 0.5 | 0.3 (−0.2) | 0.5 (0) | 0.3 | 0.2 (−0.1) | 0.3 (0) | 0.2 | 0.2 (0) | 0.2 (0) | 0.3 | 0.2 (−0.1) | 0.3 (0) |
| BC | 0.2 | 0.2 (0) | 0.2 (0) | 0.1 | 0.1 (0) | 0.1 (0) | 0.1 | 0.1 (0) | 0.1 (0) | 0.1 | 0.1 (0) | 0.1 (0) |
| $PM_{2.5}$ | 3.7 | 3.6 (−0.1) | 4.1 (+0.4) | 3.1 | 3.0 (−0.1) | 4.2 (+1.1) | 2.8 | 2.8 (0) | 3.6 (+0.8) | 3.1 | 3.1 (0) | 3.8 (+0.7) |
| | | | | | | SCB | | | | | | |
| $SO_4^{2-}$ | 1.8 | 1.8 (0) | 2.0 (+0.2) | 1.0 | 1.0 (0) | 1.3 (+0.3) | 0.9 | 1.0 (+0.1) | 1.1 (+0.2) | 1.3 | 1.2 (−0.1) | 1.5 (+0.2) |
| $NO_3^-$ | 1.2 | 1.2 (0) | 1.5 (+0.3) | 0.7 | 0.7 (0) | 1.1 (+0.4) | 0.8 | 0.8 (0) | 1.3 (+0.5) | 0.8 | 0.8 (0) | 1.3 (+0.5) |
| $NH_4^+$ | 0.6 | 0.6 (0) | 0.7 (+0.1) | 0.5 | 0.5 (0) | 0.7 (+0.2) | 0.5 | 0.5 (0) | 0.7 (+0.2) | 0.6 | 0.6 (0) | 0.7 (+0.1) |
| OC | 0.3 | 0.3 (0) | 0.4 (+0.1) | 0.2 | 0.2 (0) | 0.3 (+0.1) | 0.2 | 0.2 (0) | 0.2 (0) | 0.2 | 0.2 (0) | 0.2 (0) |
| BC | 0.2 | 0.1 (−0.1) | 0.2 (0) | 0.1 | 0.1 (0) | 0.1 (0) | 0.1 | 0.1 (0) | 0.1 (0) | 0.1 | 0.1 (0) | 0.1 (0) |
| $PM_{2.5}$ | 3.1 | 3.1 (0) | 3.7 (+0.6) | 2.0 | 2.0 (0) | 3.1 (+1.1) | 2.3 | 2.4 (0) | 3.2 (+0.9) | 2.5 | 2.5 (0) | 3.4 (+0.9) |





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**Table 3.** Simulated APDM values (%) of aerosols in ANNmet, ANNmet_ATM, and ANNall. For each season, the numbers in the brackets of column ANNmet_ATM (ANNall) were calculated as the differences in APDM values between ANNmet_ATM (ANNall) and ANNmet (ANNmet_ATM (ANNall) – ANNmet).

| Species | DJF | | | MAM | | | JJA | | | SON | | |
|---|---|---|---|---|---|---|---|---|---|---|---|---|
| | ANNmet | ANNmet_ATM | ANNall | ANNmet | ANNmet_ATM | ANNall | ANNmet | ANNmet_ATM | ANNall | ANNmet | ANNmet_ATM | ANNall |
| | | | | | | NC | | | | | | |
| $SO_4^{2-}$ | 23.6 | 23.8 (+0.2) | 24.6 (+1.0) | 17.0 | 16.7 (−0.3) | 18.2 (+1.2) | 14.7 | 14.7 (0) | 14.6 (−0.1) | 13.7 | 13.6 (−0.1) | 17.3 (+3.6) |
| $NO_3^-$ | 18.2 | 18.3 (+0.1) | 20.0 (+1.8) | 15.6 | 15.7 (0.1) | 19.2 (+3.6) | 17.0 | 17.4 (0.4) | 19.6 (+2.6) | 13.0 | 13.2 (+0.2) | 15.1 (+2.1) |
| $NH_4^+$ | 17.0 | 17.1 (+0.1) | 18.2 (+1.2) | 15.2 | 15.1 (−0.1) | 17.1 (+1.9) | 14.9 | 14.8 (−0.1) | 15.6 (+0.7) | 11.7 | 11.7 (0) | 13.4 (+1.7) |
| OC | 12.1 | 12.1 (0) | 13.1 (+1.0) | 10.7 | 8.7 (−2.0) | 11.3 (+0.6) | 8.3 | 8.2 (−0.1) | 8.5 (+0.2) | 7.7 | 7.6 (−0.1) | 8.5 (+0.8) |
| BC | 12.0 | 12.1 (+0.1) | 13.9 (+1.9) | 8.7 | 8.7 (0) | 10.4 (+1.7) | 8.2 | 8.3 (0.1) | 9.6 (+1.4) | 7.5 | 7.5 (0) | 8.9 (+1.4) |
| $PM_{2.5}$ | 16.9 | 16.9 (0) | 17.9 (+1.0) | 13.8 | 13.8 (0) | 15.8 (+2.0) | 13.7 | 13.7 (0) | 14.5 (+0.8) | 10.7 | 10.8 (+0.1) | 12.5 (+1.8) |
| | | | | | | SC | | | | | | |
| $SO_4^{2-}$ | 13.9 | 14.1 (+0.2) | 15.5 (+1.6) | 11.8 | 11.6 (−0.2) | 14.1 (+2.3) | 15.7 | 15.1 (−0.6) | 17.6 (+1.9) | 10.6 | 10.5 (−0.1) | 12.1 (+1.5) |
| $NO_3^-$ | 11.5 | 11.5 (0) | 13.0 (+1.5) | 11.7 | 11.5 (−0.2) | 18.0 (+6.3) | 17.7 | 16.7 (−1.0) | 24.5 (+6.8) | 12.2 | 11.9 (−0.3) | 16.6 (+4.4) |
| $NH_4^+$ | 9.3 | 9.2 (−0.1) | 10.9 (+1.6) | 10.5 | 10.3 (−0.2) | 14.5 (+4.0) | 16.1 | 15.4 (−0.7) | 19.8 (+3.7) | 9.5 | 9.4 (−0.1) | 13.6 (+4.1) |
| OC | 11.9 | 7.6 (−4.3) | 12.5 (+0.6) | 10.3 | 7.4 (−2.9) | 10.4 (+0.1) | 9.2 | 8.8 (−0.4) | 9.3 (+0.1) | 7.2 | 6.8 (−0.4) | 7.6 (+0.4) |
| BC | 8.6 | 7.8 (−0.8) | 10.5 (+1.9) | 7.7 | 7.2 (−0.5) | 9.8 (+2.1) | 8.3 | 8.2 (−0.1) | 10.2 (+1.9) | 6.9 | 6.9 (0) | 7.9 (+2.0) |
| $PM_{2.5}$ | 9.2 | 8.8 (−0.4) | 10.6 (+1.4) | 9.7 | 9.5 (−0.2) | 13.5 (+3.8) | 13.8 | 13.4 (−0.4) | 17.1 (+3.3) | 8.6 | 8.6 (0) | 10.7 (+2.1) |
| | | | | | | SCB | | | | | | |
| $SO_4^{2-}$ | 14.4 | 14.2 (−0.2) | 16.6 (+2.2) | 10.1 | 9.6 (−0.5) | 13.8 (+3.7) | 10.8 | 10.2 (−0.6) | 15.0 (+4.2) | 10.7 | 10.6 (−0.1) | 13.7 (+3.0) |
| $NO_3^-$ | 11.4 | 11.5 (+0.1) | 14.3 (+2.9) | 9.3 | 9.3 (0) | 15.9 (+6.6) | 12.9 | 11.3 (−1.6) | 21.6 (+8.7) | 10.5 | 9.8 (−0.7) | 14.8 (+4.3) |
| $NH_4^+$ | 9. | 9.0 (0) | 10.3 (+1.3) | 8.5 | 8.2 (−0.3) | 12.7 (+4.2) | 10.8 | 10.0 (−0.8) | 16.2 (+5.4) | 8.7 | 8.5 (−0.2) | 11.1 (+2.4) |
| OC | 7.3 | 6.8 (−0.5) | 8.9 (+1.6) | 6.6 | 5.5 (−1.1) | 8.2 (+1.6) | 5.8 | 5.8 (0) | 6.9 (+1.1) | 5.6 | 5.6 (0) | 6.2 (+0.6) |
| BC | 7.4 | 7.4 (0) | 9.5 (+2.1) | 5.3 | 5.4 (+0.1) | 8.1 (+1.8) | 5.4 | 5.4 (0) | 8.0 (+2.6) | 5.8 | 5.8 (0) | 7.7 (+1.9) |
| $PM_{2.5}$ | 8.8 | 8.8 (0) | 10.5 (+1.7) | 7.6 | 7.5 (−0.1) | 11.8 (+4.2) | 9.2 | 8.8 (−0.4) | 14.0 (+4.8) | 8.0 | 7.9 (−0.1) | 10.5 (+2.5) |





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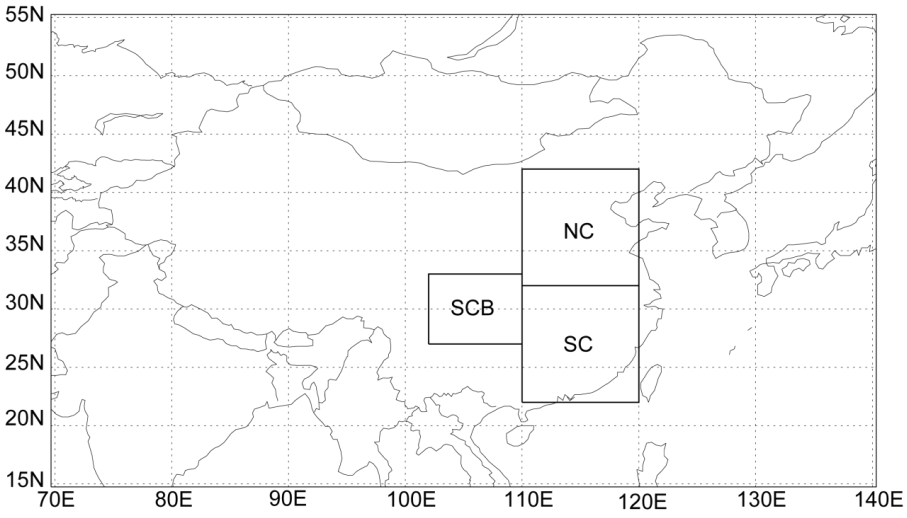


**Fig. 1.** Polluted regions examined in this study, including North China (NC, 32–42° N, 110–120° E), South China (SC, 22–32° N, 110–120° E), and the Sichuan Basin (SCB, 27–33° N, 102–110° E).





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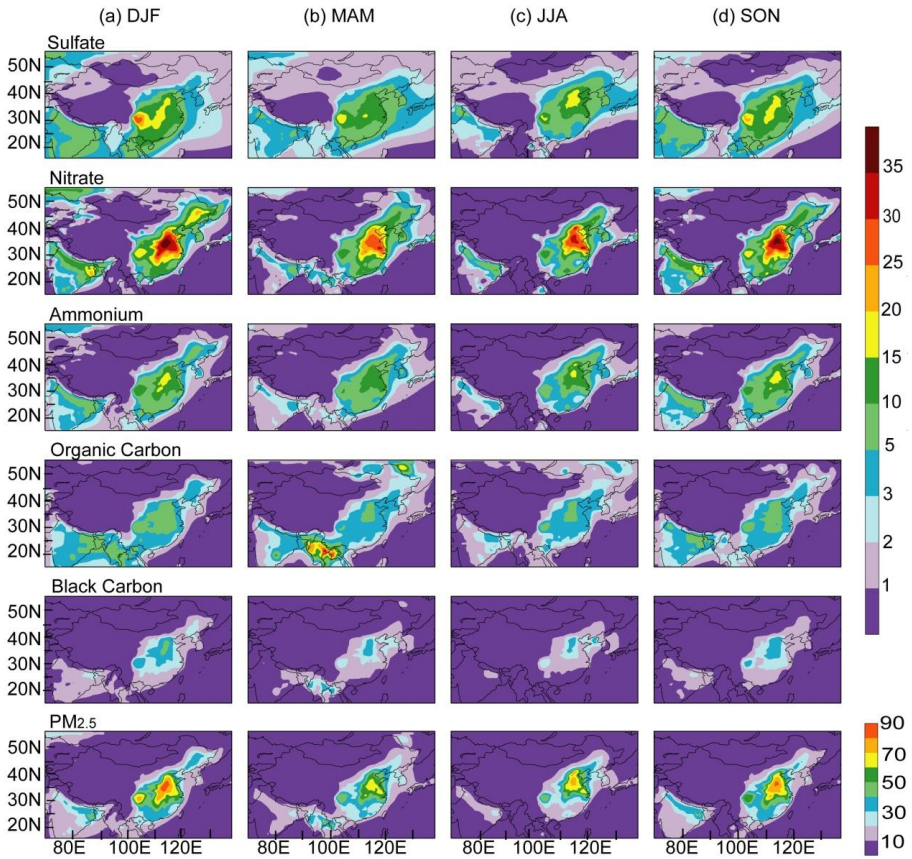


**Fig. 2.** Simulated seasonal-mean surface-layer concentrations ($\mu g\,m^{-3}$) of sulfate, nitrate, ammonium, OC, BC, and $PM_{2.5}$ in ANNmet averaged over 2004–2012.





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Simulation of the interannual variations of aerosols in China

Q. Mu and H. Liao



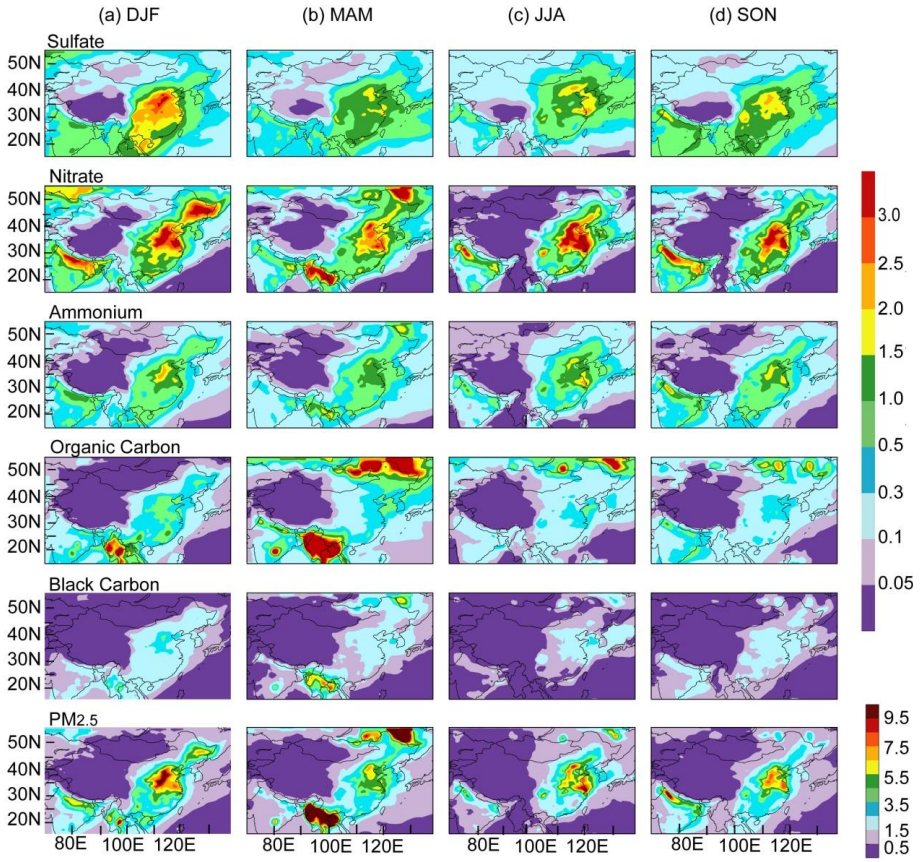


**Fig. 3.** Mean absolute deviation (MAD, $\mu g\,m^{-3}$) of surface-layer concentrations of sulfate, nitrate, ammonium, OC, BC, and $PM_{2.5}$ simulated in ANNmet for years 2004–2012.





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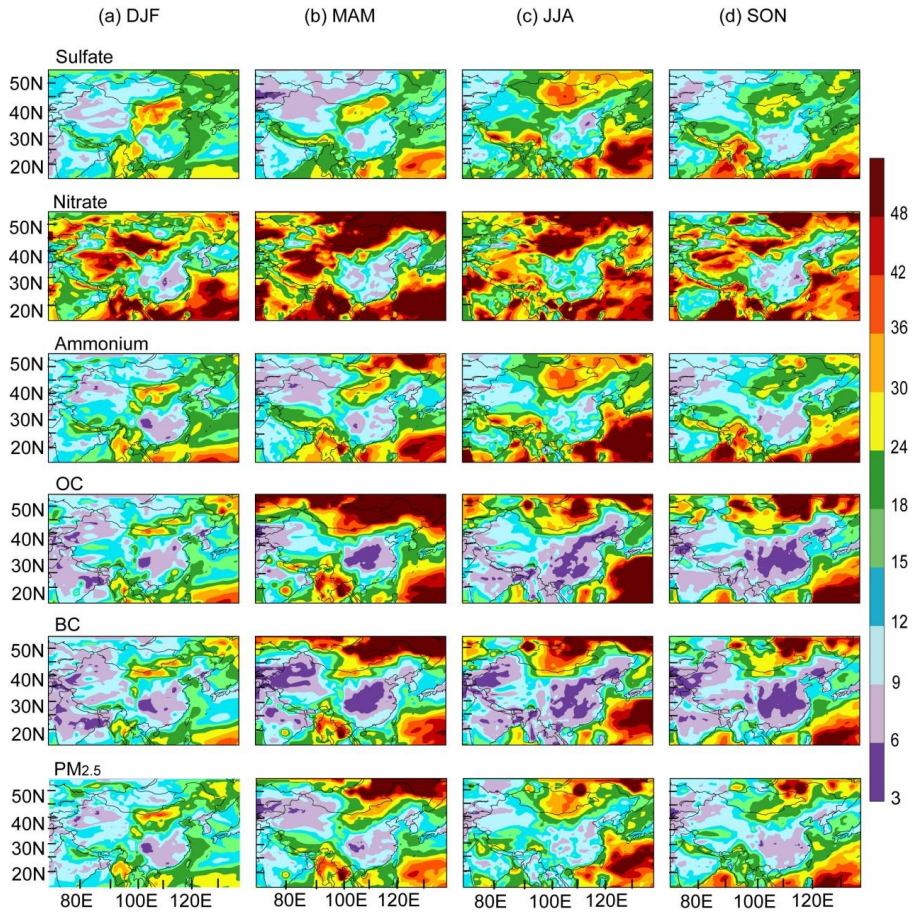


**Fig. 4.** Absolute percent departure from the mean (APDM, %) of surface-layer concentrations of sulfate, nitrate, ammonium, OC, BC, and $PM_{2.5}$ simulated in ANNmet for years 2004–2012.





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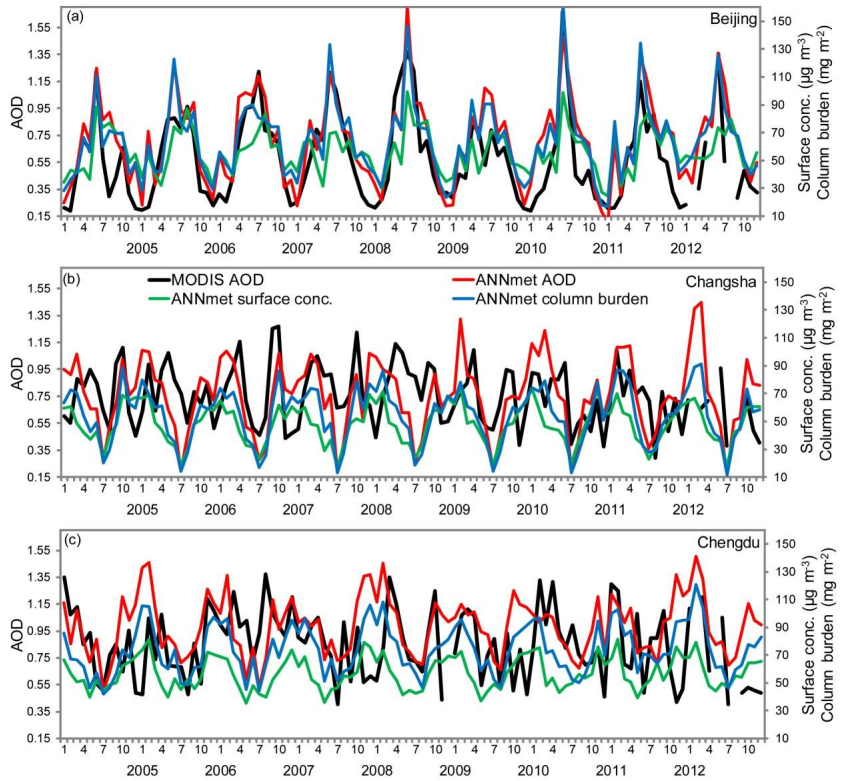


**Fig. 5.** MODIS AOD (black line, left axis), simulated AOD (red line, left axis), simulated surface-layer $PM_{2.5}$ concentrations (green line, $\mu g\,m^{-3}$, right axis), and simulated column burden of $PM_{2.5}$ (blue line, mg $m^{-2}$, right axis) over polluted cities of **(a)** Beijing, **(b)** Changsha, and **(c)** Chengdu. All simulated values were obtained from simulation ANNmet.





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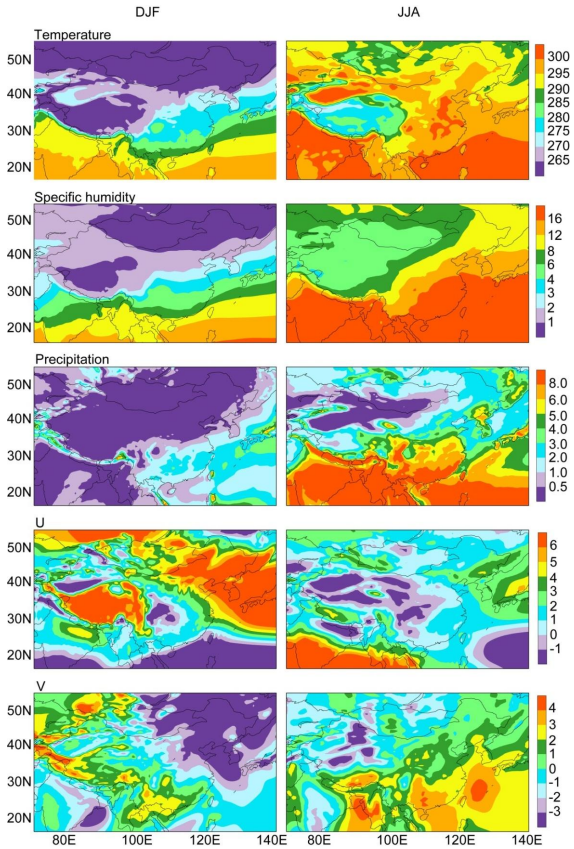


**Fig. 6.** Seasonal mean surface air temperature (K), specific humidity ($g\,kg^{-1}$), precipitation ($mm\,d^{-1}$), zonal and meridional wind at 850 hPa ($m\,s^{-1}$) in DJF and JJA. Winds that were eastward or northward were positive, and those that were westward and southward were negative. Meteorological fields were from the GEOS-5 assimilated meteorological data and were averaged over 2004–2012.





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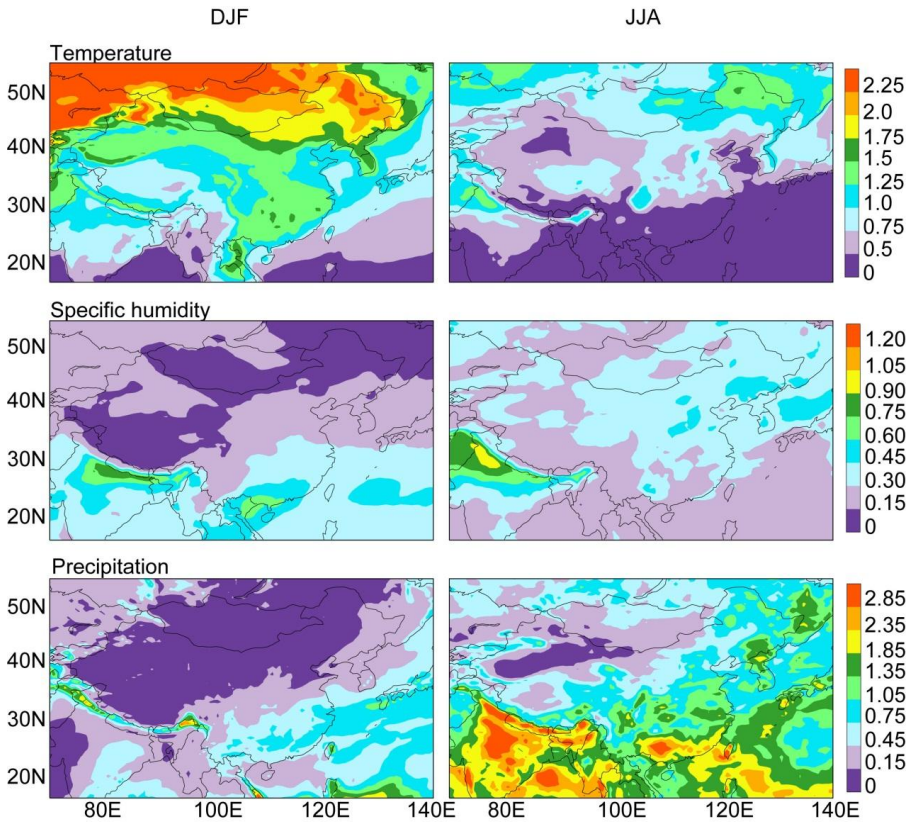


**Fig. 7.** The MAD values of surface air temperature (K), specific humidity ($g\,kg^{-1}$) and precipitation ($mm\,d^{-1}$) in DJF and JJA based on the GEOS-5 assimilated meteorological fields of 2004–2012.





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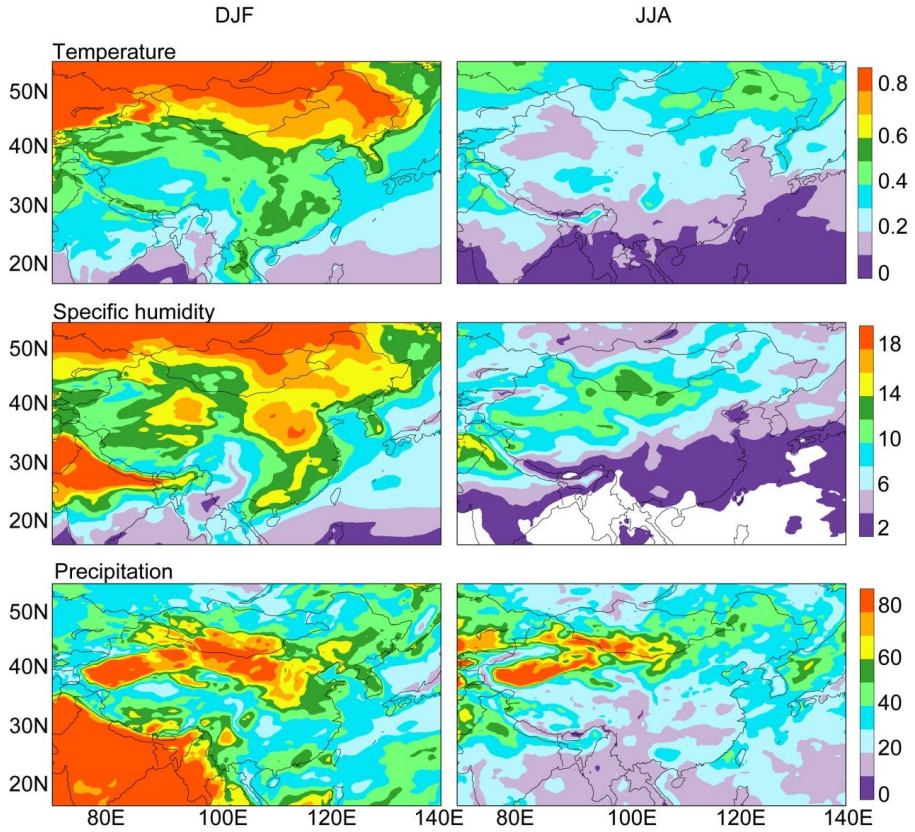


**Fig. 8.** The APDM values (%) of surface air temperature, specific humidity, and precipitation in DJF and JJA based on the GEOS-5 assimilated meteorological fields of 2004–2012.





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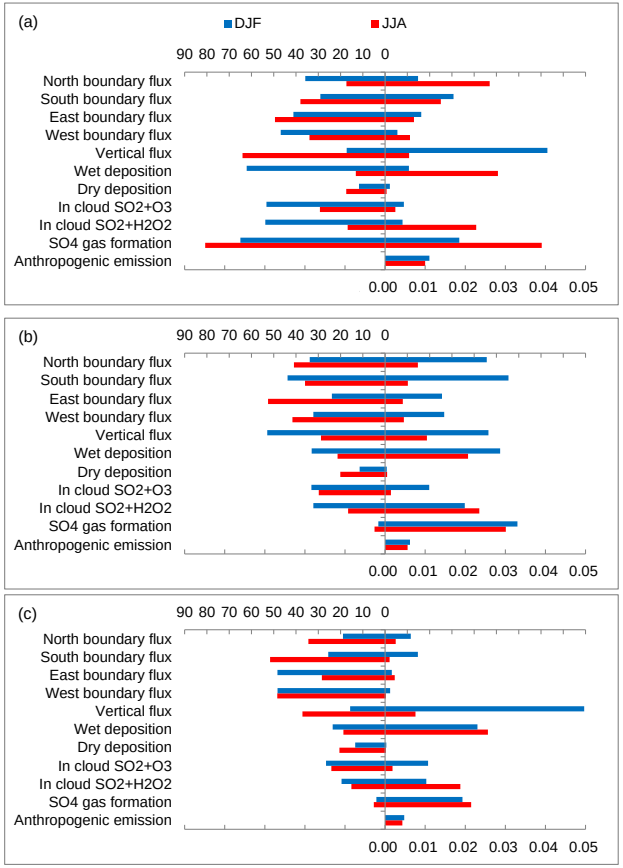


**Fig. 9.** Sulfate budget (mass flux from each process in Tg S season$^{-1}$, right) and APDM of each flux (%, left) in **(a)** NC, **(b)** SC and **(c)** SCB. Blue lines are for DJF and red lines are for JJA. Transport fluxes were calculated at the boundaries while other sources and sinks were calculated as the sums over the grid cells of the region under 1 km.





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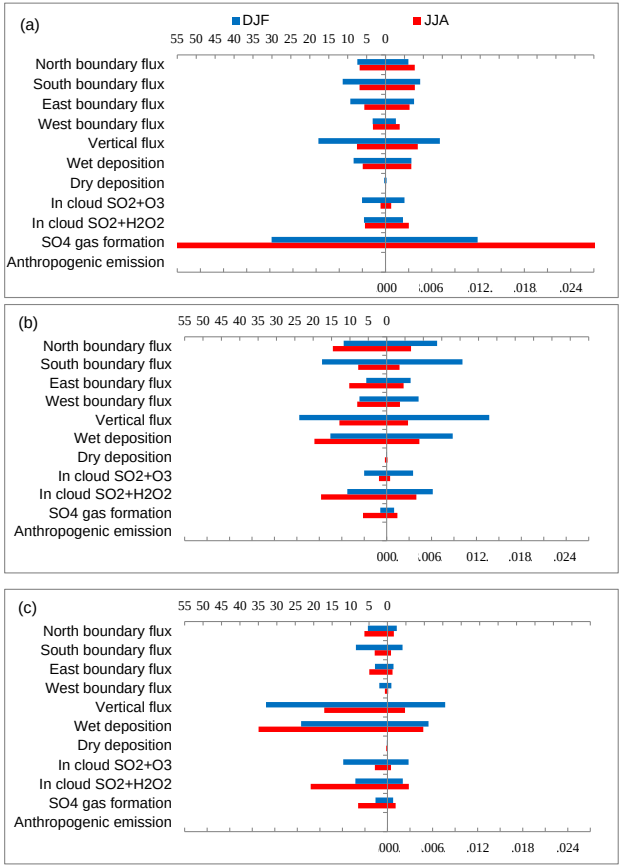


**Fig. 10.** The MAD (Tg S season$^{-1}$, right) and relative contribution (%, left) of each sulfate process (flux) in **(a)** NC, **(b)** SC, and **(c)** SCB. Blue lines are for DJF and red lines are for JJA. Transport fluxes were calculated at the boundaries while other sources and sinks were calculated as the sums over the grid cells of the region under 1 km.





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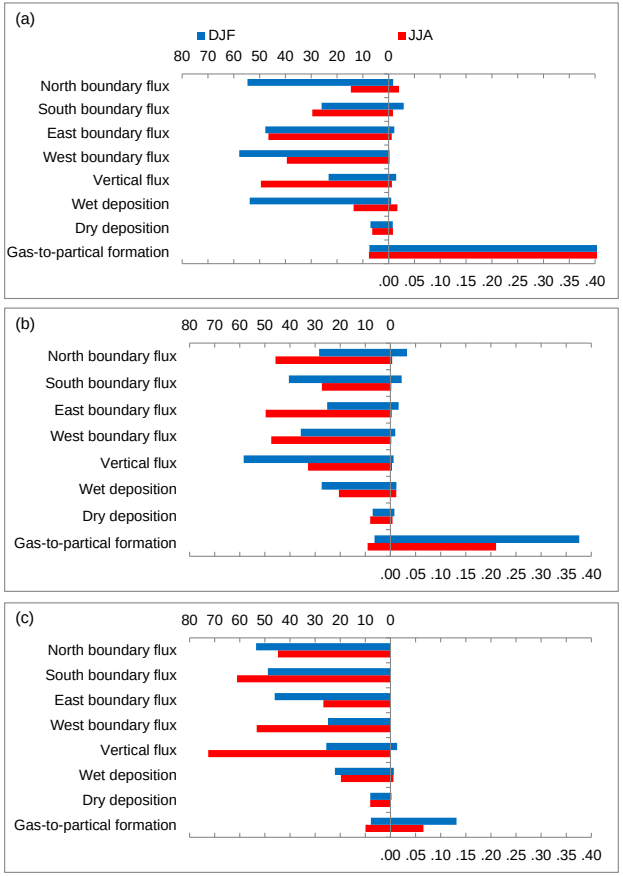


**Fig. 11.** Nitrate budget (mass flux from each process in $\mathrm{Tg\,N\,season^{-1}}$, right) and APDM of each flux (%, left) in **(a)** NC, **(b)** SC and **(c)** SCB. Blue lines are for DJF and red lines are for JJA. Transport fluxes were calculated at the boundaries while other sources and sinks were calculated as the sums over the grid cells of the region under 1 km.





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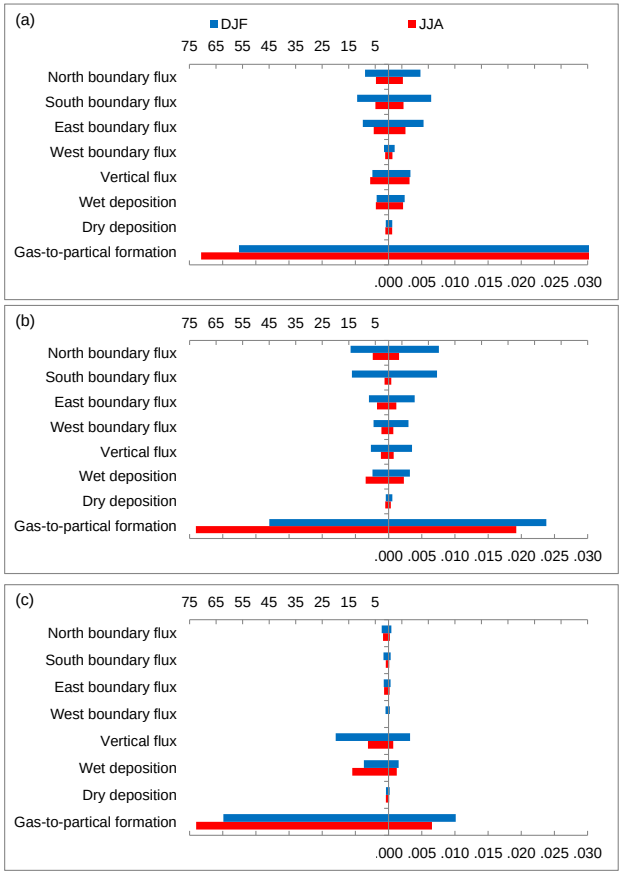


**Fig. 12.** The MAD (Tg N season$^{-1}$, right) and relative contribution (%, left) of each nitrate process (flux) in **(a)** NC, **(b)** SC, and **(c)** SCB. Blue lines are for DJF and red lines are for JJA. Transport fluxes were calculated at the boundaries while other sources and sinks were calculated as the sums over the grid cells of the region under 1 km.





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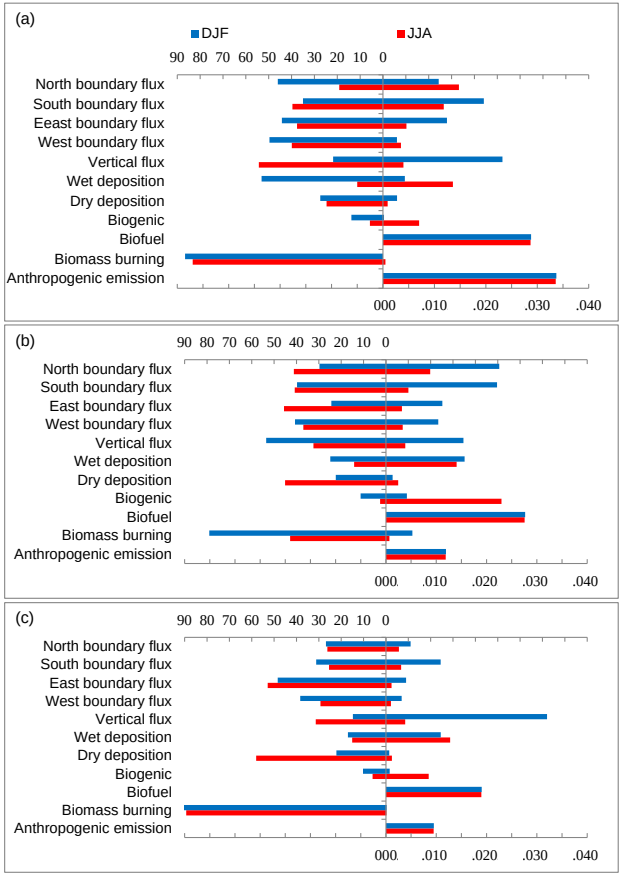


**Fig. 13.** Organic carbon budget (mass flux from each process in $TgC\,season^{-1}$, right) and APDM of each flux (%, left) in **(a)** NC, **(b)** SC and **(c)** SCB. Blue lines are for DJF and red lines are for JJA. Transport fluxes were calculated at the boundaries while other sources and sinks were calculated as the sums over the grid cells of the region under 1 km.





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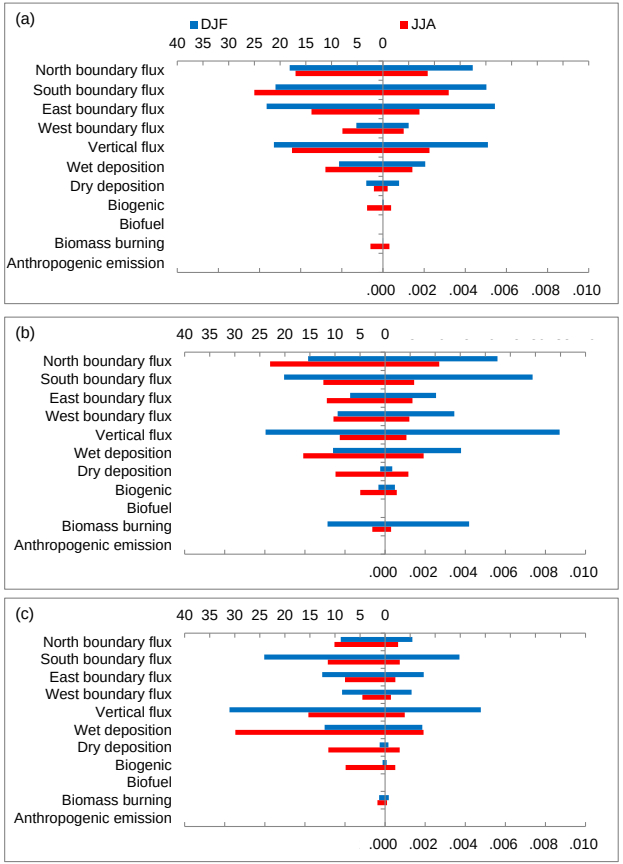


**Fig. 14.** The MAD (Tg C season$^{-1}$, right) and relative contribution (%, left) of each organic carbon process (flux) in **(a)** NC, **(b)** SC, and **(c)** SCB. Blue lines are for DJF and red lines are for JJA. Transport fluxes were calculated at the boundaries while other sources and sinks were calculated as the sums over the grid cells of the region under 1 km.





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